

# Lie Symmetry Analysis of the Charney-Hasegawa-Mima equation

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## Geophysical background:

- Charney J. G., *On the Scale of Atmospheric Motions*, Astrophysical Institute, University of Oslo (1948).
- LaCasce J. H., *Atmosphere-Ocean Dynamics*, Dept. of Geosciences, University of Oslo (2020).
- Pedlosky J., *Geophysical Fluid Dynamics*, Springer (1987).

## Methodology:

- Olver, P., *Equivalence, Invariants and Symmetry*, Cambridge University Press (1995).

## Lie Symmetries of the CHM equation for the special case $\beta = F = 1$ :

- Hounkonnou M. N., Kabir M. M., *Hasegawa - Mima - Charney - Obukhov Equation: Symmetry Reductions and Solutions*, Int. J. Contemp. Math. Sciences, Vol.3, p. 145 - 157 (2008).

The Charney-Hasegawa-Mima equation:

$$\frac{\partial}{\partial t}(\Delta u - Fu) + \beta \frac{\partial u}{\partial x} + [u, \Delta u] = 0,$$

- $u = u(t, x, y)$  ... stream function  
 $t$  ... temporal coordinate  
 $(x, y)$  ... spatial coordinates  
 $\beta > 0, F \geq 0$  ... constants

$$\beta = \beta_0 \frac{L^2}{U}, \quad \beta_0 = \frac{2\Omega \cos \theta}{R_E}, \quad F = \left(\frac{L}{R}\right)^2,$$

$$R = \frac{\sqrt{gD}}{f} - \text{Rossby radius of deformation}$$

$\Omega$  – Earth's rotation rate,  $\theta$  – latitude,  $R_E$  – Earth's radius

$$D/L/U - \text{scales} \sim (10\text{km}/1000\text{km}/10\text{m} \cdot \text{s}^{-1})$$

The Coriolis parameter  $f = 2\Omega \sin \theta$  - influence of the Coriolis force on the fluid.  $f(\theta) \approx f(\theta_0) + \beta_0 y$ , where  $y \equiv R_E(\theta - \theta_0)$ .

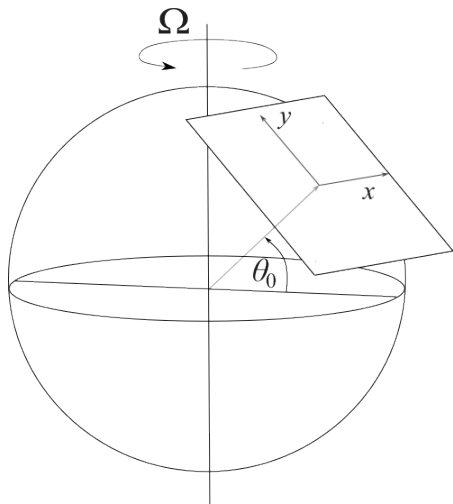


Figure:  $\beta$  - plane model of the CHM equation

Assumptions under which the CHM equation holds and was derived:

- The flow is non-viscous and the density  $\rho$  of the fluid is constant.

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- Hydrostatic balance between pressure gradient and gravity

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- Geostrophic balance between horizontal pressure gradient and Coriolis pressure.
- Small dimensionless Rossby and temporal Rossby numbers

$$R_o = \frac{U}{fL}, \quad R_T = \frac{1}{fT}, \quad U/T - \text{horizontal velocity/time scale.}$$

# Infinitesimal method

General  $n$ -th order system of PDEs:

$$\Delta_{\kappa}(\mathbf{x}, \mathbf{u}^{(n)}) = 0, \quad \kappa = 1, \dots, m,$$

where  $(\mathbf{x}, \mathbf{u}^{(n)}) \in J^n Y$ .

Identification with variety

$$\mathcal{S}_{\Delta} = \{(\mathbf{x}, \mathbf{u}^{(n)}) \mid \Delta_{\kappa}(\mathbf{x}, \mathbf{u}^{(n)}) = 0, \quad \kappa = 1, \dots, m, \}$$

- Solution - any function  $s(\mathbf{x})$ , such that its graph of  $n$ -th prolongation lies in  $\mathcal{S}_{\Delta}$  (i.e.  $(\mathbf{x}, s^{(n)}(\mathbf{x})) \subseteq \mathcal{S}_{\Delta}$ ).
- The Lie point symmetry - any smooth point transformation (action of the Lie group  $G$ ), which maps smooth solutions to smooth solutions.

$\implies \mathcal{S}_{\Delta}$  is  $G$  - invariant.

## Infinitesimal method

Lie algebra  $\mathfrak{g}$  associated to Lie group  $G$  is spanned by infinitesimal generators

$$X = \sum_{i=1}^p \xi^i(\mathbf{x}, \mathbf{u}) \frac{\partial}{\partial x^i} + \sum_{\alpha=1}^q \phi^\alpha(\mathbf{x}, \mathbf{u}) \frac{\partial}{\partial u^\alpha}. \quad (1)$$

Then  $n$ -th prolongation of the vector field  $X$  has a form

$$X^{(n)} = \sum_{i=1}^p \xi^i(\mathbf{x}, \mathbf{u}) \frac{\partial}{\partial x^i} + \sum_{\alpha=1}^q \sum_{\#J=j=0}^n \phi_J^\alpha(\mathbf{x}, \mathbf{u}^{(j)}) \frac{\partial}{\partial u_J^\alpha}, \quad (2)$$

with coefficients

$$\phi_J^\alpha = D_J Q^\alpha + \sum_{i=1}^p \xi^i u_{J,i}^\alpha, \quad \alpha = 1, \dots, q, \quad (3)$$

where  $D_J Q^\alpha$  is the total derivative of the characteristic function

$$Q^\alpha = \phi^\alpha(\mathbf{x}, \mathbf{u}) - \sum_{i=1}^p \xi^i(\mathbf{x}, \mathbf{u}) \frac{\partial u^\alpha}{\partial x^i}, \quad \alpha = 1, \dots, q. \quad (4)$$

# Infinitesimal symmetry criterion

## Theorem

A connected Lie group  $G$  is a symmetry group of the fully regular system of differential equations  $\Delta = 0$ , if and only if the classical infinitesimal symmetry conditions

$$X^{(n)}(\Delta_\kappa) = 0, \quad \kappa = 1, \dots, r, \quad \text{whenever } \Delta = 0,$$

hold for every infinitesimal generator  $X \in \mathfrak{g}$  of  $G$ .

## Algorithm

- Set symmetry criterion for given PDE or system of PDEs

$$X^{(n)}(\Delta_\kappa) = 0, \quad \kappa = 1, \dots, r, \quad \text{whenever } \Delta = 0.$$

- Derive coefficients  $\phi_J^\alpha$ .
- Solve the system of *determining equations*.

The CHM equation with applied Laplacian and Jacobian:

$$u_{txx} + u_{tyy} - Fu_t + \beta u_x + u_x u_{xxy} + u_x u_{yyy} - u_y u_{xxx} - u_y u_{xyy} = 0.$$

Infinitesimal generator  $X(t, x, y, u)$  over  $J^0Y$

$$X(t, x, y, u) = \tau \frac{\partial}{\partial t} + \chi \frac{\partial}{\partial x} + \psi \frac{\partial}{\partial y} + \mu \frac{\partial}{\partial u}.$$

The third prolongation  $X^{(3)}(t, x, y, u)$  over  $J^3Y$

$$\begin{aligned} X^{(3)} = & \tau \frac{\partial}{\partial t} + \chi \frac{\partial}{\partial x} + \psi \frac{\partial}{\partial y} + \mu \frac{\partial}{\partial u} + \mu^t \frac{\partial}{\partial u_t} + \mu^x \frac{\partial}{\partial u_x} + \mu^y \frac{\partial}{\partial u_y} + \\ & + \mu^{tt} \frac{\partial}{\partial u_{tt}} + \mu^{tx} \frac{\partial}{\partial u_{tx}} + \dots + \mu^{xyy} \frac{\partial}{\partial u_{xyy}} + \mu^{yyy} \frac{\partial}{\partial u_{yyy}}. \end{aligned}$$

Infinitesimal symmetry criterion:

$$X^{(3)}(\Delta_1) = \mu^{txx} + \mu^{tyy} - F\mu^t + \beta\mu^x + \mu^x(u_{xxy} + u_{yyy}) + \\ + u_x(\mu^{xxy} + \mu^{yyy}) - \mu^y(u_{xxx} + u_{xyy}) - u_y(\mu^{xxx} + \mu^{xyy}) = 0, \quad (5)$$

whenever

$$u_{txx} + u_{tyy} + Fu_t - \beta u_x - u_x u_{xxy} - u_x u_{yyy} + u_y u_{xxx} + u_y u_{xyy} = 0. \quad (6)$$

We derive the coefficients  $\mu^t, \mu^x, \dots, \mu^{yyy}$ , for example

$$\mu^t = D_t \mu - (D_t \tau) u_t - (D_t \chi) u_x - (D_t \psi) u_y, \\ \mu^{xxy} = D_{xxy} \mu - D_{xxy}(\tau u_t) - D_{xxy}(\chi u_x) - D_{xxy}(\psi u_y) + \\ + \tau u_{txxy} + \chi u_{xxxy} + \psi u_{xxyy},$$

a insert them into equation (5), together with expressed term from eq. (6)

$$u_{txx} = -u_{tyy} + Fu_t - \beta u_x - u_x u_{xxy} - u_x u_{yyy} + u_y u_{xxx} + u_y u_{xyy}.$$

We obtain the polynomial with indeterminates  $u, u_t, u_x, \dots, u_{yyy}$ :

$$\begin{aligned}
 & 1 \cdot (\mu_{txx} + \mu_{tyy} - F\mu_t + \beta\mu_x) + \\
 & u_t \cdot (\mu_{xxu} + \mu_{yyu} - \beta\tau_x - 2F\chi_x) + \\
 & u_x \cdot (2\mu_{txu} + \mu_{xxy} + \mu_{yyy} - \chi_{txx} - \chi_{tyy} + F\chi_t + \beta\chi_x + \beta\tau_t) + \\
 & u_y \cdot (2\mu_{tyu} - \mu_{xxx} - \mu_{xyy} - \psi_{txx} - \psi_{tyy} + F\psi_t - \beta\psi_x) + \\
 & u_t u_x \cdot (2\mu_{xuu}) + \\
 & \quad + \\
 & \quad \vdots \\
 & \quad + \\
 & u_y u_{xyy} \cdot (3\psi_y - \tau_t - \chi_x - \mu_u) + \\
 & u_y u_{yyy} \cdot (-\psi_x + \psi_x) = 0.
 \end{aligned}$$

All brackets has to be equal to zero  $\implies$  linear PDEs with constant coef.



Solution of linear PDEs with constant coefficients:

- Macaulay2 – Linux software for computations in commutative algebra and algebraic geometry.
- We solve in Macaulay2 those equations without general coefficients  $\beta, F$ .
- Remaining equations with general coefficients  $\beta, F$  can be solved easily by hand.

Algorithm core in Macaulay2:

- Linear partial differential equations with constant coefficients have a same structure as vectors of polynomials

$$2\chi_{xy} + 3\chi_{xx} + \psi_{yy} - 2\mu_{yu} \mapsto \begin{pmatrix} 0 \\ 2\partial_x\partial_y + 3(\partial_x)^2 \\ (\partial_y)^2 \\ -2\partial_y\partial_u \end{pmatrix},$$

- *Fundamental Principle of Ehrenpreis–Palamodov.*

System of linear PDEs has different solution depending on the  $F$  value

$$F = \left( \frac{L}{R} \right)^2 = \frac{f^2 L^2}{gD},$$

where  $L$  – characteristic length and  $R$  – Rossby radius of deformation.

- 1  $F \approx 0$  (near equator) – Rossby radius of deformation is large. Planet's rotation has dominant effect on the flow development – zonal flows, jet streams, ocean currents.
- 2  $F > 0$  (mid-latitudes) – Influences of the planet's rotation and changes in the flow (caused by pressure/temperature gradient, topography etc.) are in balance.
- 3  $F \gg 0$  (near north/south poles) – Rossby radius of deformation is small, planet's rotation has neglectable effect on the fluid flow.

Solution with assumption  $F > 0$  (balance, mid-latitudes):

$$\tau(t) = c_1 + c_5 t,$$

$$\chi(t, x, y) = c_2 - \frac{\beta}{F} c_5 t - c_6 y,$$

$$\psi(t, x, y) = c_3 + \frac{\beta}{F} c_6 t + c_6 x,$$

$$\mu(t, x, y, u) = c_4 + \frac{\beta^2}{F^2} c_6 t + \frac{\beta}{F} c_6 x + \frac{\beta}{F} c_5 y - c_5 u.$$

Infinitesimal generators of Lie algebra  $\mathfrak{g}_{sym}$ :

$$X_1 = \frac{\partial}{\partial t}, \quad X_2 = \frac{\partial}{\partial x}, \quad X_3 = \frac{\partial}{\partial y}, \quad X_4 = \frac{\partial}{\partial u},$$

$$X_5 = Ft \frac{\partial}{\partial t} - \beta t \frac{\partial}{\partial x} + (\beta y - Fu) \frac{\partial}{\partial u},$$

$$X_6 = -F^2 y \frac{\partial}{\partial x} + (\beta Ft + F^2 x) \frac{\partial}{\partial y} + (\beta^2 t + \beta F x) \frac{\partial}{\partial u}.$$

Corresponding Lie point transformations ( $\alpha_i$  – group parameter):

- Generators  $X_1, X_2, X_3, X_4$ :

$$(\tilde{t}, \tilde{x}, \tilde{y}, \tilde{u}) = (t + \alpha_1, x + \alpha_2, y + \alpha_3, u + \alpha_4), \quad \alpha_1, \alpha_2, \alpha_3, \alpha_4 \in \mathbb{R}.$$

- Generator  $X_5$

$$\tilde{t} = te^{F\alpha_5},$$

$$\tilde{x} = x + \frac{\beta t}{F}(1 - e^{F\alpha_5}),$$

$$\tilde{y} = y,$$

$$\tilde{u} = \frac{\beta y}{F}(1 - e^{-F\alpha_5}) + ue^{-F\alpha_5}.$$

- Generator  $X_6$ , where  $K = \frac{\beta}{F}$

$$\tilde{t} = t,$$

$$\tilde{x} = x \cos(F^2 \alpha_6) - y \sin(F^2 \alpha_6) + Kt(\cos(F^2 \alpha_6) - 1),$$

$$\tilde{y} = x \sin(F^2 \alpha_6) + y \cos(F^2 \alpha_6) + Kt \sin(F^2 \alpha_6),$$

$$\tilde{u} = Kx \sin(F^2 \alpha_6) + Ky (\cos(F^2 \alpha_6) - 1) + K^2 t \sin(F^2 \alpha_6) + u.$$

Abbreviating  $F^2$  and reorganizing terms

$$\tilde{t} = t,$$

$$\tilde{x} = (x + Kt) \cos(\alpha_6) - y \sin(\alpha_6) - Kt,$$

$$\tilde{y} = (x + Kt) \sin(\alpha_6) + y \cos(\alpha_6),$$

$$\tilde{u} = K(x + Kt) \sin(\alpha_6) + Ky (\cos(\alpha_6) - 1) + u.$$

# Geometrical interpretation

- Generators  $X_1, X_2, X_3, X_4$  – translation in coordinates  $t, x, y, u$ .

$$(\tilde{t}, \tilde{x}, \tilde{y}, \tilde{u}) = (t + \alpha_1, x + \alpha_2, y + \alpha_3, u + \alpha_4), \quad \alpha_1, \alpha_2, \alpha_3, \alpha_4 \in \mathbb{R}.$$

– doesn't have impact on the fluid velocity fields

- Generator  $X_5$  – time-length contraction/extension in  $(t, x)$  plane

$$\tilde{t} = te^{F\alpha_5}, \quad \tilde{x} = x + \frac{\beta t}{F}(1 - e^{F\alpha_5})$$

$$\implies \tilde{t}_2 - \tilde{t}_1 = (t_2 - t_1)e^{F\alpha_5}$$

- $\alpha_5 > 0$  – time extension
- $\alpha_5 < 0$  – time contraction

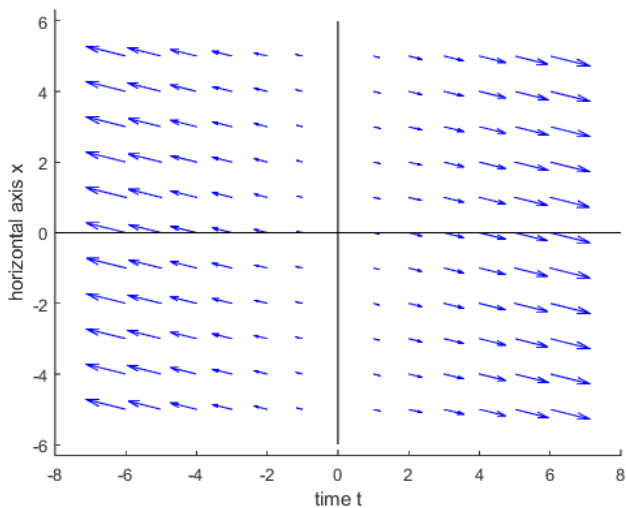


Figure: Vector field  $X_5$ ,  $\beta = 0.25$ ,  $F = 1$

Length/time change ratio: - phase velocity  $K = \frac{\beta}{F}$  is preserved ( $x$  dir.)

$$\frac{\tilde{x} - x}{\tilde{t} - t} = -\frac{\beta}{F}, \quad \tilde{x} = x - \frac{\beta}{F}(\tilde{t} - t).$$



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1) Example of stream function

$$u = y \implies \tilde{u} = \tilde{y} \left( e^{-F\alpha_6} + \frac{\beta}{F} (1 - e^{-F\alpha_6}) \right)$$

Change in flow velocity fields:

$$\left( \frac{\partial u}{\partial y}, -\frac{\partial u}{\partial x} \right) = (1, 0) \implies \left( \frac{\partial \tilde{u}}{\partial \tilde{y}}, -\frac{\partial \tilde{u}}{\partial \tilde{x}} \right) = \left( e^{-F\alpha_6} + \frac{\beta}{F} (1 - e^{-F\alpha_6}), 0 \right)$$

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$$\left( \frac{\partial u}{\partial y}, -\frac{\partial u}{\partial x} \right) = (1, 0) \implies \left( \frac{\partial \tilde{u}}{\partial \tilde{y}}, -\frac{\partial \tilde{u}}{\partial \tilde{x}} \right) = \left( e^{-F\alpha_6} + \frac{\beta}{F} (1 - e^{-F\alpha_6}), 0 \right)$$

2) Example of stream function

$$u = x \implies \tilde{u} = \frac{\beta}{F} \tilde{y} (1 - e^{-F\alpha_6}) + \tilde{x} e^{-F\alpha_6}$$

Change in flow velocity fields:

$$\left( \frac{\partial u}{\partial y}, -\frac{\partial u}{\partial x} \right) = (0, -1) \implies \left( \frac{\partial \tilde{u}}{\partial \tilde{y}}, -\frac{\partial \tilde{u}}{\partial \tilde{x}} \right) = \left( \frac{\beta}{F} (1 - e^{-F\alpha_6}), e^{-F\alpha_6} \right)$$

- Generator  $X_6$  – rotation around the origin with respect to the phase velocity in  $x$  direction  $K = \frac{\beta}{F}$ .

$$\tilde{t} = t,$$

$$\tilde{x} = (x + Kt) \cos(\alpha_6) - y \sin(\alpha_6) - Kt,$$

$$\tilde{y} = (x + Kt) \sin(\alpha_6) + y \cos(\alpha_6),$$

$$\tilde{u} = K(x + Kt) \sin(\alpha_6) + Ky (\cos(\alpha_6) - 1) + u.$$

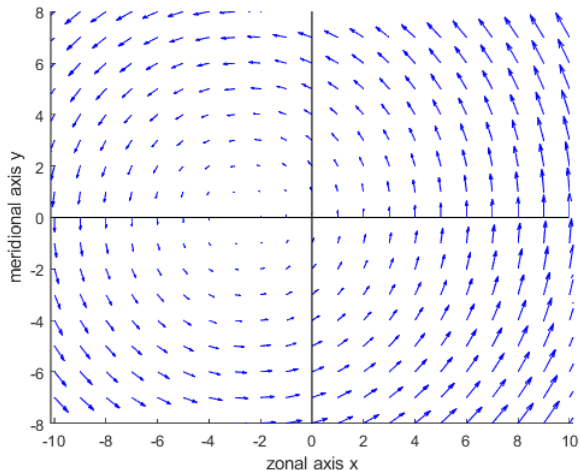


Figure: Vector field  $X_6$ ,  $\beta = 0.25$ ,  $F = 1$ ,  $t = 10$

Solution with assumption  $F \gg 0$  (neglectable rotation influence):

$$\tau(t) = c_1 + c_5 t,$$

$$\chi(y) = c_2 - c_6 y,$$

$$\psi(x) = c_3 + c_6 x,$$

$$\mu(u) = c_4 - c_5 u.$$

Infinitesimal generators of Lie algebra  $\mathfrak{g}_{sym}$ :

$$X_1 = \frac{\partial}{\partial t}, \quad X_2 = \frac{\partial}{\partial x}, \quad X_3 = \frac{\partial}{\partial y}, \quad X_4 = \frac{\partial}{\partial u},$$

$$X_5 = t \frac{\partial}{\partial t} - u \frac{\partial}{\partial u},$$

$$X_6 = -y \frac{\partial}{\partial x} + x \frac{\partial}{\partial y}.$$

Corresponding Lie-point transformations and interpretation

- Generators  $X_1, X_2, X_3, X_4$  – translations
- Generator  $X_5$  – scaling in temporary coordinate  $t$

$$(\tilde{t}, \tilde{x}, \tilde{y}, \tilde{u}) = (te^{\alpha_5}, x, y, ue^{-\alpha_5}).$$

- Generator  $X_6$  – rotation in  $\beta$ -plane

$$(\tilde{t}, \tilde{x}, \tilde{y}, \tilde{u}) = (t, x \cos(\alpha_6), y \sin(\alpha_6), u).$$

Solution with assumption  $F \approx 0$  (dominant rotation influence):

$$\tau(t) = c_1,$$

$$\chi(t, x, y) = -h(t),$$

$$\psi(t, x, y) = c_2,$$

$$\mu(t, x, y, u) = f(t) + h'(t)y.$$

Infinitesimal generators of Lie algebra  $\mathfrak{g}_{sym}$ :

$$X_1 = \frac{\partial}{\partial t}, \quad X_2 = \frac{\partial}{\partial y},$$

$$X_3 = -h(t)\frac{\partial}{\partial x} + h'(t)y\frac{\partial}{\partial u}, \quad X_4 = f(t)\frac{\partial}{\partial u}.$$

## Corresponding Lie-point transformations and interpretation

- Generators  $X_1, X_2$  – translations in  $t, y$  coordinates
- Generator  $X_3$  – time-dependent motion along  $x$ -coordinate  
– by setting  $\alpha_3 = t$ ,  $h(t)$  represents planet's peripheral velocity (zonal direction)

$$\tilde{t} = t,$$

$$\tilde{x} = -h(t)\alpha_3 + x,$$

$$\tilde{y} = y,$$

$$\tilde{u} = h'(t)y\alpha_3 + u.$$

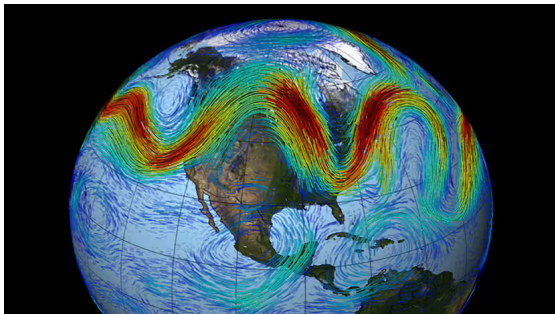
- Generator  $X_4$  – any time-dependent change in stream function values

$$(\tilde{t}, \tilde{x}, \tilde{y}, \tilde{u}) = (t, x, y, f(t)\alpha_4 + u).$$



## Conclusion

- For reasonably large  $F > 0$  (mid-latitudes), the CHM equation has Lie symmetries of translations, time-scaling and rotation, which preserve phase velocity  $\frac{\beta}{F}$  in  $x$ -direction.
- Special cases of  $F$  value (rather theoretical meaning):
  - ①  $F \approx 0$  (planet's rotation is dominant, equator)  $\implies$  time-dependent motion symmetries along  $x$  and  $u$  coordinates
  - ②  $F \gg 0$  (planet's rotation is neglectable, poles)  $\implies$  rotation symmetry in spatial  $(x, y)$  coordinates and scaling in temporary  $t$  coordinate



## Conclusion

- Translation symmetry is present in all  $F$  instances.
- The most general solution of the CHM equation (for particular  $F$  instance) – application of all  $G_{sym}$  transformations to coordinates of the known solution  $u = u(t, x, y)$ .

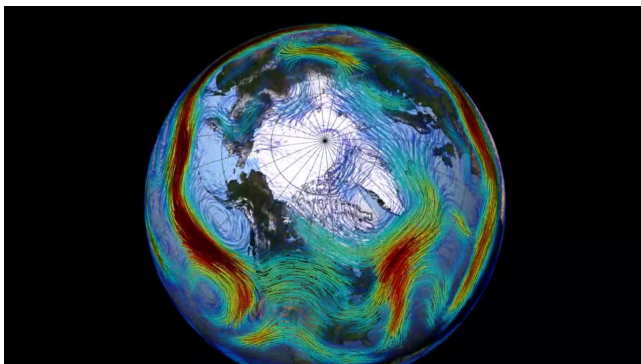


Figure: Rossby waves